

# ASSESSING THE PREDICTABILITY OF TOTAL SEDIMENT TRANSPORT RATE FOR UNIT STREAM POWER APPROACH

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**Abstract**— Sediment transport is a significant factor in effective design of hydraulic structures and open channel water conveyance. Estimation of the amount of sediment material which a specific flow can carry is one of the major issues of sedimentation research. Unit stream power is defined as the time rate of potential energy per unit weight of water, and is found to be the dominant factor in the determination of total sediment concentration being carried by the (Yang, 1972). Yang's (1979) unit stream power approach for total load sediment transport function is applied to test the predictability of total sediment transport rate by using flume data sets of Samaga et al (1986), Willis et al (1972) and river data set of Middle Loup river of Hubbel et al (1959). The performance of the Yang's sediment transport function for different data sets has been evaluated by finding statistical parameters, such as root mean square error (RMSE), discrepancy ratio and inequality coefficient. From the evaluation, it is observed that the Yang (1979) total load function under predicts the sediment transport rate for the three mixtures  $M_1$ ,  $M_2$ ,  $M_3$  and predicts well for mixture  $M_4$  of Samaga et.al data set giving good result compared to other mixtures. For both Willis et al (1972) data set and the river data of Hubbel et al (1959) total load function of Yang (1979), both under predicts as well as over predicts sometimes. Yang total load equation gives better results for samaga et al (1986) data sets for the four mixture.

**Index Terms**— Fall velocity, incipient motion, sediment transport, shear velocity, total sediment load, unit stream power and Yang.C.T.

## 1 INTRODUCTION

The subject sediment transport and flow in alluvial stream are gaining importance with the utilization of water resources. Considerable development has taken place in the field of fluvial hydraulics which is considered as complicated branch of engineering.

Estimation of the amount of sediment material, which a specific flow can carry, is one of the major issues of sedimentation research. In alluvial river system, river banks will erode, sediment will be deposited and flood plains and side channels will undergo modification with time. Effect of sedimentation in river reduces carrying capacity which may lead to flood water damage to surrounding area.

Contributors to the development of total load sediment transport theories refers to the names of Laursen (1958), Garde and Albertson (1961), Ackers and White (1973), Engelund and Hanson (1967), Yang (1973, 1979, 1984) etc. The relationship between rate of sediment transport and rate of potential energy expenditure has been studied in detail. The concept of the rate of work done should be related to the rate of energy expenditure was used by Bagnold (1966). It was demonstrated by Yang (1972) that the rate of sediment transport depends on the unit stream power more than any other hydraulic parameter. Unit stream power, defined as the time rate of potential energy expenditure per unit weight of water is shown to be the dominant factor in the determination of total sediment concentration. Yang's (1973) unit stream power equation for the computation of total sediment concentration includes the incipient motion criteria while Yang's (1979) unit stream power

equation for the computation of total sediment concentration is obtained without using any criteria for incipient motion. Due to the uncertainties involved in determining the flow conditions precisely at incipient motion, the present paper will examine the possibility of developing an accurate unit stream power equation for total load or total bed material load in the sand size range without using any criteria for incipient motion.

## 2 METHODOLOGY

Yang (1979) approach is used in the present analysis to predictability the total load transport rate.

### 2.1 Chih Ted Yang (1979)

The concept was introduced and developed by Bagnold (1966) of stream power. The unit stream power of Yang (1972) can be derived from the stream power. This slightly different theory relates the rate of potential energy dissipation of the unit weight of water. Chih Ted Yang (1979), suggest that the following relation provides the best correlation between total sediment concentration  $C_t$  and unit stream power ( $V_s$ ).

$$\log C_t = I_1 + J_1 \log \left( \frac{V_s}{\omega} \right) \quad (1)$$

Further from multiple regression analysis with 463 sets of laboratory data from uniform flows, the following equations were obtained for  $I_1$  and  $J_1$  parameters.

$$I_1 = 5.165 - 0.153 \log \left( \frac{\omega d_s}{v} \right) - 0.297 \log \left( \frac{u_*}{\omega} \right) \quad (2)$$

$$J_1 = 1.78 - 0.36 \log \left( \frac{\omega d_s}{v} \right) - 0.48 \log \left( \frac{u_*}{\omega} \right) \quad (3)$$

Eqs. 1, 2 and 3 were proposed by Yang (1979) for the prediction of total sediment concentration  $C_t$  in parts per million by weight for the particles in the sand size range with  $d_s = d_{50}$  = median sieve diameter of bed material.

### 2.1 Step by step computational procedure for Yang

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(1979).

- Input parameter:  $Q, D, V, S, g, v, d_{50}$ .
- Output :  $Q_t$ .

**Computational procedure :**

The procedure for the computation is summarized as follows:

1. Compute the value of overall shear velocity  $U_*$  from

$$u_* = (gDS)^{0.5} \tag{4}$$

2. Calculate fall velocity of sediment particles  $\omega$  using  $d_s = d_{50}$

$$\omega = \left(\frac{10 v}{d_s}\right) \left\{ \left[ 1 + \left(\frac{0.01 \Delta g d_s^3}{v}\right) \right]^{0.5} - 1 \right\} \tag{5}$$

3. Calculate the factor  $I_1$  from Eq., 2.
4. Calculate the factor  $J_1$  from Eq., 3.
5. Determine the total sediment transport rate  $Q_t$  in ppm by weight from Eq., 1.

**3 DATA COLLECTION**

In the present study following flume and field data sets have been used for testing Yang (1979) approach.

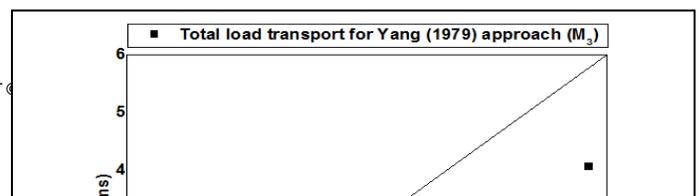
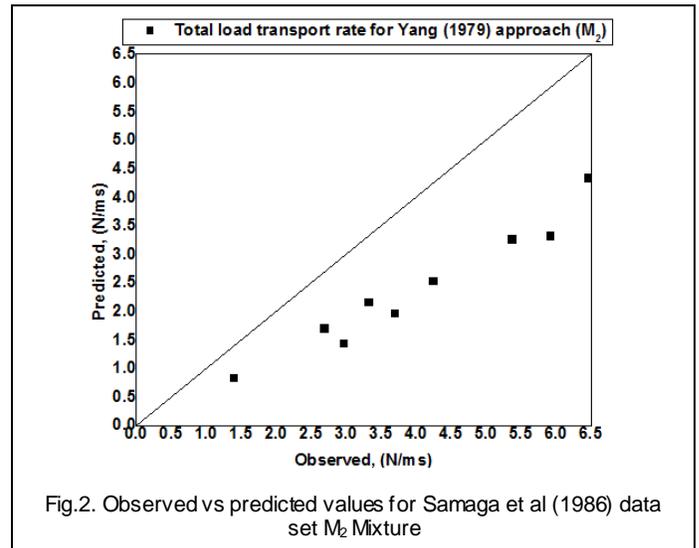
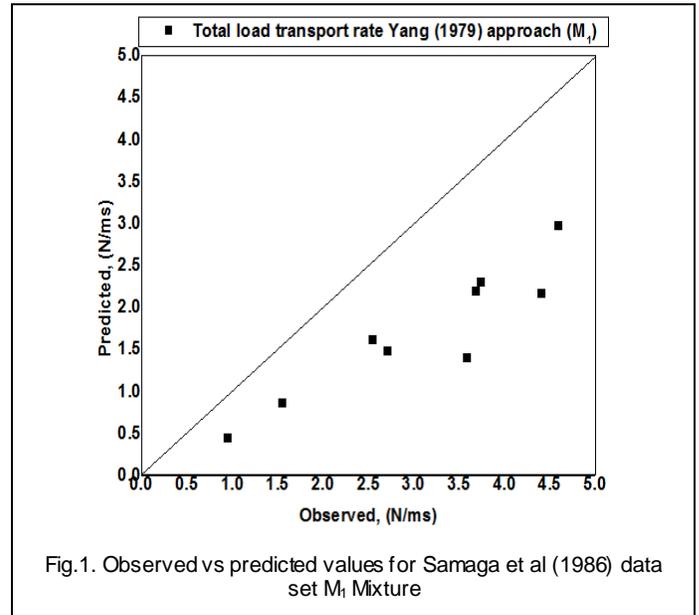
1. Samaga et al flume data (1986).
2. Willis et al flume data (1972).
3. Middle loup river Hubbel et al (1959).

Total load transport rate using Yang (1979) relation are predicted for the four mixtures ( $M_1, M_2, M_3, M_4$ ) having different arithmetic mean diameter and standard deviation as given in Samaga et al (1986) flume data set. The value of grain size  $d_{50}$  for the four sediment mixtures are  $M_1 = 0.350$  mm,  $M_2 = 0.290$  mm,  $M_3 = 0.270$  mm and  $M_4 = 0.200$  mm and for Willis et al (1972) grain size  $d_{50} = 0.1$  mm median sieve diameter. However Yang (1979) total load function is also predicted using field data of Middle loup river Hubbel et al (1968). The computed total load transport using flume data are plotted in Fig. 1, 2, 3, 4, 5 and for field data plotted in Fig. 6. Observed and predicted values are plotted in origin to check the predictability of Yang (1979) total load function. Solid line represents the line of equality. Values above the solid line represents values are over prediction and values below the solid line represents the value are under prediction of total load.

**4 RESULT ANALYSIS**

From the Fig. 1, it can be observed that the sediment transport rate for mixture  $M_1$  values are below the line of equality so the observed and predicted are plotted in origin graph represent the value are underpredict. While mixture  $M_2$  and  $M_3$  also below the line of equality so the observed and predicted are plotted in origin graph represent the value are underpredict. For the mixture  $M_4$  it could be concluded that the shows values are nearer the line of equality so the observed and predicted values are plotted in origin graph represent the good result for  $M_4$  mixture. For laboratory data willis et al (1972) observed vs predicted graph shown in Fig. 5 values are above as well as below

the line of equality. So, the observed and predicted values are plotted in origin graph is over predict as well as under predict and for field data Middle loup river Hubbel et al (1959) origin graph shown in Fig 6 values are above as well as below the line of equality. So the observed and predicted values are plotted in origin graph is over predict as well as under predict. Comparison between observed values and predicted values obtained for Samaga et al (1986) data set is shown in table 1, 2, 3 and 4.



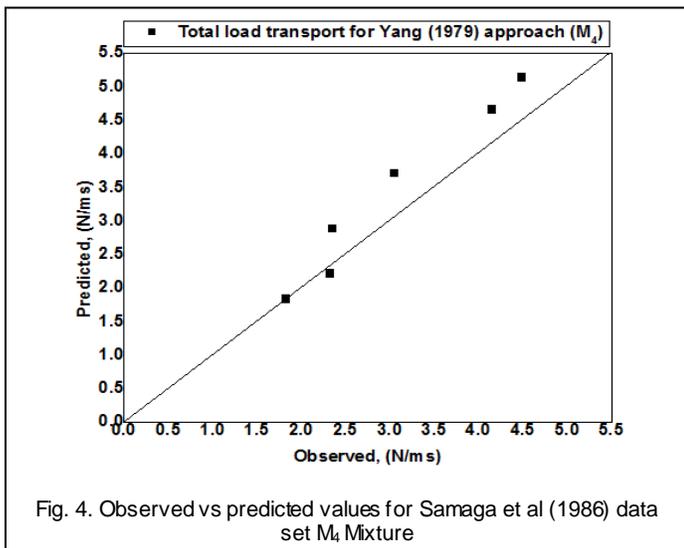


TABLE 1  
COMPARISON BETWEEN PREDICTED VALUES AND OBSERVED VALUES FOR SAMAGA ET AL, DATA SET-M<sub>1</sub>

Yang Total load transport rate (N/ms)			
M <sub>1</sub> Mixture			
Sr. No	Observed Q <sub>t</sub> (N/ms)	Predicted Q <sub>t</sub> (N/ms)	(Observed - Predicted)*100/Predicted
1	1.545	0.866663404	-43.90528127
2	2.546	1.628463012	-36.03837345
3	3.737	2.309682088	-38.19421761
4	2.705	1.488652206	-44.96664672
5	3.68	2.196103668	-40.3232699
6	4.587	2.982457338	-34.98021936
7	0.943	0.454169146	-51.83784244
8	3.582	1.408523867	-60.67772567
9	4.399	2.169838236	-50.67428424
			<b>-44.62198452</b>

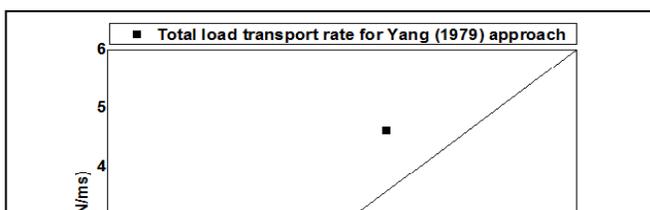
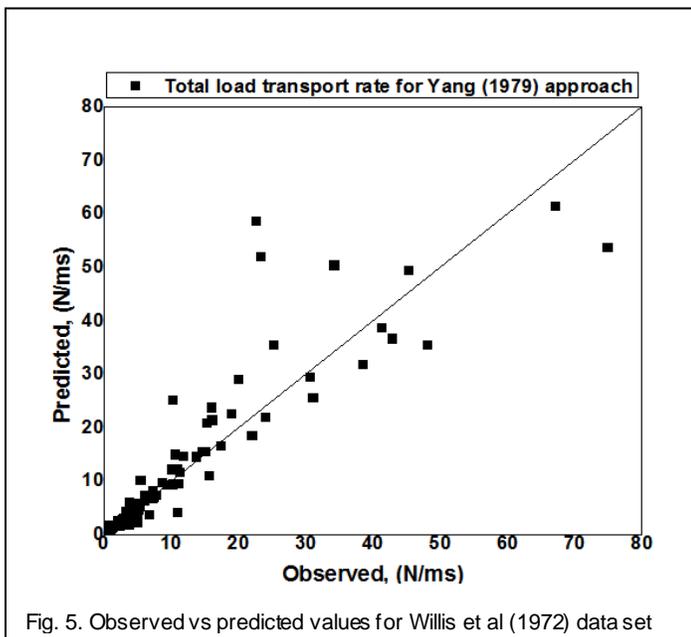


TABLE 2  
COMPARISON BETWEEN PREDICTED VALUES AND OBSERVED VALUES FOR SAMAGA ET AL, DATA SET-M<sub>2</sub>

Yang Total load transport rate (N/ms)			
M <sub>2</sub> Mixture			
Sr. No	Observed Q <sub>t</sub> (N/ms)	Predicted Q <sub>t</sub> (N/ms)	(Observed - Predicted)*100/Predicted
1	1.389	0.847280884	-39.00065627
2	2.682	1.716133094	-36.0129346
3	3.689	1.973399739	-46.50583522
4	2.956	1.451595706	-50.89324403
5	4.24	2.541656644	-40.05526784
6	5.913	3.330805883	-43.66978044
7	3.32	2.174101583	-34.51501256
8	5.369	3.268806947	-39.11702465
9	6.462	4.348968642	-32.69934011
			<b>-40.27434397</b>

TABLE 3  
COMPARISON BETWEEN PREDICTED VALUES AND OBSERVED VALUES FOR SAMAGA ET AL, DATA SET-M<sub>3</sub>

Yang Total load transport rate (N/ms)			
M <sub>3</sub> Mixture			
Sr. No	Observed Q <sub>t</sub> (N/ms)	Predicted Q <sub>t</sub> (N/ms)	(Observed - Predicted)*100/Predicted
1	5.51	3.1367993	-43.07079344
2	5.765	4.0844178	-29.15146979
3	3.61	2.2516402	-37.62769518
4	2.718	1.2372069	-54.48098051
5	3.512	1.9038214	-45.79096361
6	3.852	3.0936835	-19.68630587
7	3.001	1.6758522	-44.15687548
8	1.897	1.0947261	-42.29171736
9	4.918	3.471243	-29.41758746
			<b>-38.40826541</b>

TABLE 4  
COMPARISON BETWEEN PREDICTED VALUES AND OBSERVED VALUES FOR SAMAGA ET AL, DATA SET-M<sub>4</sub>

Yang Total load transport rate (N/ms)			
M <sub>4</sub> Mixture			
Sr. No	Observed Q <sub>t</sub> (N/ms)	Predicted Q <sub>t</sub> (N/ms)	(Observed - Predicted)*100/Predicted
1	2.34	2.8956627	23.74627103
2	4.476	5.1463373	14.9762585
3	2.315	2.2236182	-3.947378787
4	4.143	4.6685543	12.68535621
5	1.82	1.8449986	1.373546746
6	3.043	3.7184669	22.19740124
			<b>11.83857582</b>

## 5 STATICAL PARAMETER TO CHECK THE PREDICTABILITY OF YANG (1979)

Various statistical parameters such as Root Mean Square Error (RMSE), Discrepancy Ratio (r) and Inequality Coefficient (U) are calculated to analyse the predicted results of the Yang's (1979) macroscopic approach by using Samaga et al (1986) flume data for the four sediment mixture (M<sub>1</sub>, M<sub>2</sub>, M<sub>3</sub>, M<sub>4</sub>), flume data of Willis et al (1972) and field data of Middle loup river Hubbel et al (1968). Summary of calculated statistical parameters for total load transport rate are shown in table 5, 6 and 7 respectively.

TABLE 5  
SUMMARY OF STATISTICAL PARAMETER OBTAINED USING YANG'S 1979 TOTAL LOAD FUNCTION FOR SAMAGA ET AL. (1986)

Statistical Parameters					
Sr. No.	Type of mixture	RMSE	Discrepancy ratio, r	Inequality coefficient	Avg. % error(MNE)
1	M-1	1.47299	0.55378	0.284841	-44.6220
2	M-2	1.70534	0.59725	0.246796	-40.2743
3	M-3	1.49623	0.61591	0.223736	-38.4082
4	M-4	0.499925	1.11838	0.073485	11.8386

**TABLE 6**  
SUMMARY OF STATISTICAL PARAMETER OBTAINED USING 1979 TOTAL LOAD FUNCTION FOR WILLIS ET AL.(1972)

Statistical Parameter				
Sr. No.	RMSE	Discrepancy ratio, r	Inequality coefficient	Avg. % error(MNE)
1	6.4142	0.1757	1.210027	21.0027

**TABLE 7**  
SUMMARY OF STATISTICAL PARAMETER OBTAINED USING YANG'S 1979 TOTAL LOAD FUNCTION FOR MIDDLE LOUP RIVER HUBBEL ET AL.(1959)

Statistical Parameter				
Sr. No.	RMSE	Discrepancy ratio, r	Inequality coefficient	Avg. % error(MNE)
1	1.80401	0.623242	0.388431	-37.6758

## 6 CONCLUSION

Analysis of predicted results and observed values for the Yang's total load function can be summarized with the following findings:

- It was observed that the predicted results of total load transport rate for  $M_1$ ,  $M_2$  and  $M_3$  mixture is good giving nearly equal values in Yang (1979) approach of Samaga et al (1986) data
- It was also observed that Yang's approach under predicts total load transport rate for all the three mixture  $M_1$ ,  $M_2$  and  $M_3$  of Samaga et al (1986) data.
- It was also observed that the value of Yang (1979) approach for total load transport for  $M_4$  mixture provide good results as compared to the mixtures  $M_1$ ,  $M_2$  and  $M_3$ .
- Yang (1979) total load function over predict as well as underpredict for Willis et al (1972) flume data set and Middle loup river Hubbel et al (1970) field data set.

## NOTATIONS

TABLE  
NOTATION

A	=	Area (m <sup>2</sup> )
Q	=	Discharge (m <sup>3</sup> /s)
W	=	Width of flume (m)
d	=	Diameter of particles (mm)
D	=	Water depth (m)
R	=	Hydraulic radius
S	=	Slope (m/m)
U	=	Average velocity (m/s)
U*	=	Shear velocity

$\tau$	=	Bed shear stress
V	=	Average water velocity (m/s)
$V_{cr}$	=	Critical average water velocity
$\omega$	=	Fall velocity
g	=	Acceleration due to gravity
$\nu$	=	Kinematic viscosity
Ct	=	Total Concentration in parts per million by weight

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